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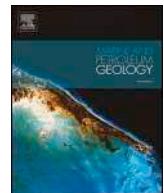
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Research paper

Kinematic analyses of brittle roto-translational planar and listric faults based on various rotational to translational velocities of the faulted blocks

Soumyajit Mukherjee^{a,*}, Lokesh Tayade^b^a Department of Earth Sciences, Indian Institute of Technology Bombay, Powai, Mumbai, 400 076, Maharashtra, India^b Department of Earth and Climate Science, Indian Institute of Science Education and Research Pune, Dr. Homi Bhabha Road, Pashan, Pune, 411 008, Maharashtra, India

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ABSTRACT

Research on rotational/roto-translational faults have been relatively less than translational faults. This is despite perfect translational faults seldom occur in mega-scale. This work models slickenside lineation patterns on planar and listric fault planes for different ratios of rates of rotation to the uniform speed towards a specific geographic direction (“ ω/v ”) of faulted blocks. Curves simulated for the planar fault planes are fitted with natural examples of such roto-translational faults. In our 10 chosen examples, the “ ω/v ” ratio ranges between 0.2 (fault at San Miguelito range, Central México) and as high as 2.4 (fault at Hyogoken-Nambu, Japan). Knowing the “ ω/v ” ratios from terrains can better interpret its tectonics, For example, if “ v ” and the total time of duration of the deformation are known from some other studies, even the total amount of rotation can be estimated. The far-reaching implication of this study is that, as several hydrocarbon reserves and terrains with landslides are related to roto-translational faults, this work will be a stepping-stone in petroleum geosciences (e.g., stability of hydrocarbon reserve) and engineering geology (e.g., stability of slope).

1. Introduction

Deciphering kinematics of faulting is one of the fundamental exercises in structural geology (e.g., Rowland and Sibson, 2001). Having diverse geometries, slickensides may develop on fault planes indicating either 180° apart two possible directions or a unique direction of slip (Doblas et al., 1997; Doblas, 1998; Mukherjee, 2014, 2015; Dasgupta and Mukherjee, 2017).

In nature/ in meso-scale/ in field, purely translational faults (Fig. 1a) with perfectly straight slickensides are found rarely. A true rotational fault (Fig. 1.b1) has also not yet been reported in the geoscience literature. Natural slickensides, on the other hand, are usually gently to moderately curved. This indicates that the faulted blocks rotate to some amount when they translate (review in Mukherjee and Khonsari, 2017). Such lineations indicate neither a first translation followed by rotation nor a vice versa, as in those cases two distinct lineation patterns would have developed one perfectly linear and another perfectly circular. Only in cases of ideal dip slip and strike slip book-shelf gliding (Mukherjee, 2018; Mukherjee and Khonsari, 2018), significantly rotated crustal blocks can leave straight slickensides at the fault surface i.e., at the interfaces between the two books/crustal blocks.

Mandal and Chakraborty (1989) refer the natural faults with

simultaneous rotation and translation as the roto-translational faults (Fig. 1c). Rep. Fig. 1 presents the global distribution of such faults where curved slickensides have been noted/photographed by the previous authors. Few papers on theoretical analyses of rotational fault kinematics based on tangent lineation diagrams using curved slickenlines are available (e.g., Twiss and Gefell, 1990; Twiss et al., 1991). Roto-translational faults can be planar or listric (referred in Marshak and Mitra, 1988; Feng et al., 2016a,b) and can be numerous in specific terrains (Nevin, 1949).

Listric faults, on the other hand, have classically been presented implicitly in structural geology text mostly to be dip-slip normal faults (e.g., Fossen 2016; Blue Mountain Fault, Australia: Mackenzier (1993); field and numerical models: Hjelle 2014). However, I. strike-slip and II. oblique-slip listric faults have also been reported cursorily in the literature. For example, I. strike-slip listric faults have been reported in mega-scale from both the continental and oceanic crusts: at the releasing bend at the Loreto basin (Dorsey et al., 1995), ridge basin (Baratoux et al., 2011; Feng et al., 2018a), southern California (May et al., 1993), Loma Prieta bend in the San Andreas Fault (Dietz and Ellsworth, 1990; Tearpock and Bischke, 2002), the pull-apart Bhima basin (south India) (Dey, 2015), the boundary between Eurasian continental crust and Pacific oceanic crust (Taiwan Petroleum Exploration Division Chinese Petroleum Corporation, 1992), the Brevard fault zone

* Corresponding author.

E-mail address: soumyajitm@gmail.com (S. Mukherjee).

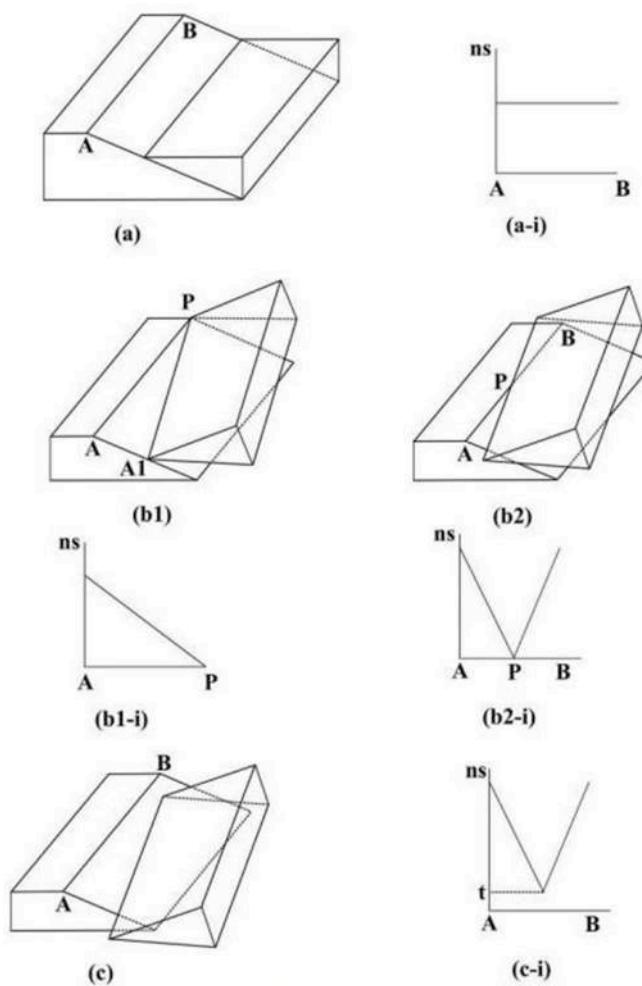


Fig. 1. Ideal faults. **a.** Translational normal fault. **a-i.** Net-slip (ns) is constant. **b1-b2.** Pure rotational fault. **b1-i.** Net-slip is zero at P and highest at A. **b2-i.** Net-slip is zero at p and highest at A and B. **c.** Roto-translational fault. **c-i.** Its net-slip profile. Reproduced from Fig. 1 of Mukherjee and Khonsari (2017).

(southern Appalachians) (Vauchez, 1987), to name a few. Rodgers and Chinnery (1969) describe strain and displacement patterns for strike-slip faults but do not simulate stretching lineation geometries on them. II. Oblique slip listric faults (Rep. Fig. 2) have been briefly reported from the North Minwun Basin (Myanmar) (Morley, 2017), the eastern margin of the Dagg ridge (New Zealand) (Keary et al., 2013). Seminoe Mountans (U.S.) (Bergh and Snock, 1992), the San Andreas fault zone (Seeger and Armbruster, 1995; Rust, 1998), syn-sedimentary slide from the Waitemata basin (New Zealand) (Sporri and Rowland, 2007), half-graben structures (Jayko and Bursik, 2012), the 12-May-2008 Sichuan seismicity in China (Zhang et al., 2010: high-angle reverse fault), the pop-up structures (Nabavi et al., 2017) etc. Unlike these authors, Claypool et al. (2002) describe a palm tree-like network of oblique-slip listric faults from Fiordland (New Zealand).

Besides merely reported (Marshak and Mitra, 1988), quantification of curvature of listric faults does not exist in the structural geological and the tectonic literature till date. Davis et al. (2019) discuss recently the kinematics of listric faults in 3D. While they describe how shear fractures are associated with such faults, the lineation pattern developed on the listric faults has not been explained.

Rotational/roto-translational faults are associated with

hydrocarbon reserves (Michelsen and Andersen, 1982; Afifi et al., 2016) and landslides (e.g., Geist, 2000). Stability of structural hydrocarbon reserves depends on the faults that define them (Neil et al., 1993). Stability of a slope would depend presumably on the “ ω/v ” ratio of the fault that initiates its slide. Therefore, understanding the kinematics of such faults will be of great importance.

This article has two parts. A. Kinematic analysis of roto-translational faults with planar and curvi-planar/listric fault surfaces. This is achieved by modeling the loci of slickenside developed on such planar fault planes for specific ratios of rotation rate to the translation rate (ω/v) of the faulted blocks. The simulated curves are then correlated with the field-photographs of curved slickensides. Field works by the first author over last 20 years did not find good examples of curved slickensides from Indian terrains. So we chose previous authors' published photographs for the present work. B. For different ratios of “ ω/v ”, geometries of slickensides on curved fault planes are also simulated. Since curved lineations on curved fault planes have merely been reported, but no images are available, the curved that we simulate in ideal cases remain uncorrelated with the natural examples.

2. Curved slickensides and fault block rotation (rates)

Twiss and Gefell (1990) describe basic kinematics of curved slickenfibres (a kind of slickenlines). However, they do not estimate the rotation rate with respect to the translational velocity. Kusky et al. (1997) report up to 55° of curvature of the lineation on fault plane. A review of literature indicates that co-seismic faulting can produce curved slickenfibres (e.g., Fig. 12 in Liu-Zeng et al., 2010), but exceptions do exist (Fig. 6b in Pan et al., 2014). Otsubo et al. (2013) report the most strikingly curved slickenside on fault plane related to the 11-April-2011 Mw = 6.8 earthquake from the Iwaki city (Japan). On the other hand, Pan et al. (2014) report a gently curved co-seismic striations on fault planes (their Fig. 6b) as a result of 12-May-2008 Mw = 7.9 seismicity in Wenchun (China).

Slip rates for translational faults of various types have been deduced, which usually range from a few mm- (Papanikolaou et al., 2005: using paleoseismologic throw-rate data from trenches and scarp profiles) up to a few cm per year (Kohn et al., 2004: based on *in-situ* Th-Pb dating of monazite). Mukherjee and Khonsari (2018) literature survey in their Table 1 reports a maximum rotation rate of crustal block $40^{\circ}\text{ Ma}^{-1}$ from North Iceland volcanic rift zone. Paleomagnetic studies have revealed considerable rotation and rotation rates from several terrains (review in Table 1). However, from these places, curved slickensides have not been reported. Price and Scott (1994) provide a much reduced rate of rotation of 3° Ma^{-1} from SW Turkey, a different terrain, due to inherent local tectonics. Horst et al. (2018) recently report a pronounced rotation of the crustal block up to 105° from the Húsavík-Flatey Fault (Iceland). Nur et al. (1989) described the tectonic reasons for significant variation of rotation rate of crustal blocks in different parts of the Earth.

3. Model

3.1. Planar fault planes

We model the geometries of curved slickensides developed at different ratios of angular (ω) to linear velocities (v) of the faulted blocks using the programming software RStudio (Version 1.0.153-2009-20017). While the angular velocity represents the rotational component of faulting, the linear velocity denotes that for the translational component of movement between the two faulted blocks. The modeled curves generated were then fitted with already published 10 good quality field photographs of planar fault planes with prominent curved

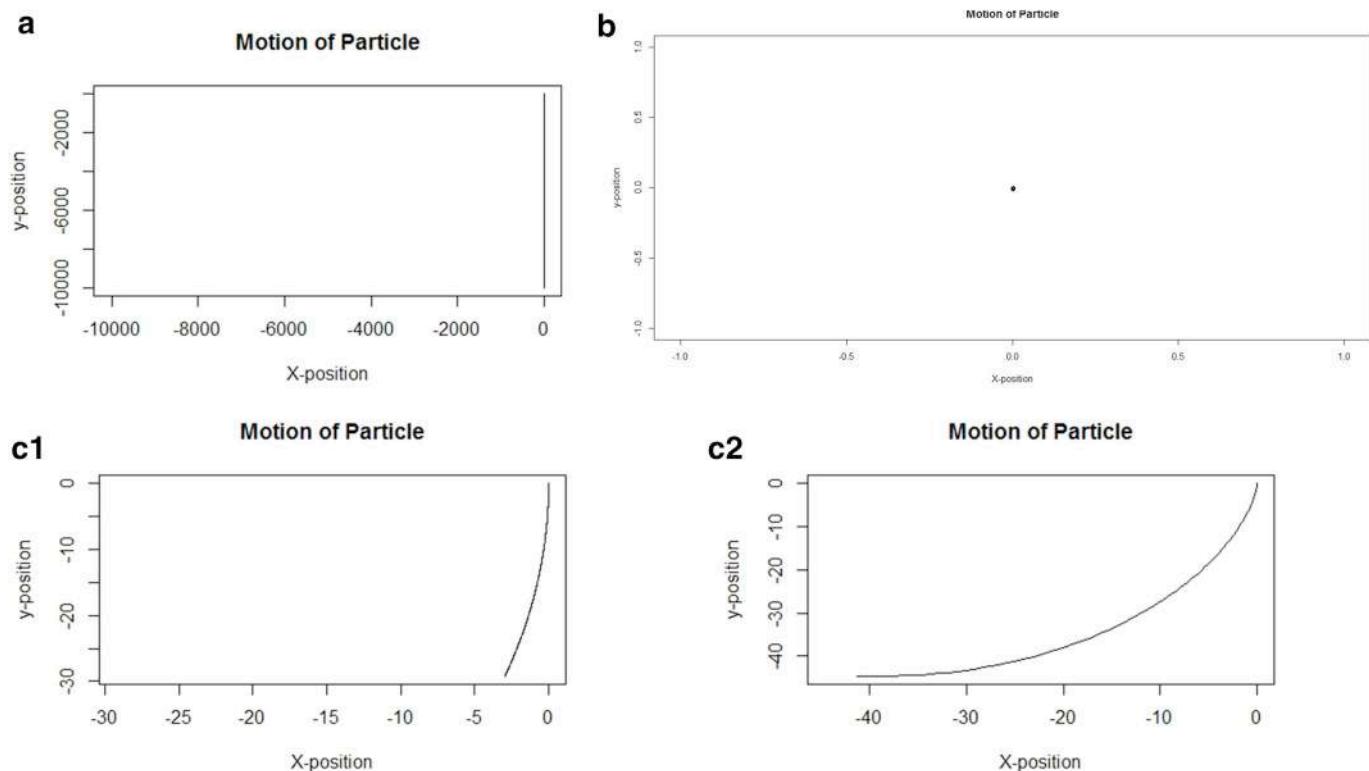


Fig. 2. The modeled lineation geometries on planar fault planes in various simple cases. **a.** $\omega = 0$ case, for translational faults such as Fig. 1a. **b.** $v = 0$ case, for rotational faults such as Fig. 1b1-i. **c1.** $\omega \neq 0, v \neq 0$ case (here $\omega/v = 1.3$), for roto-translational faults, such as Fig. 1c. **c2.** $\omega \neq 0, v \neq 0$ case (here $\omega/v = 5.6$), for roto-translational faults, such as Fig. 1c.

(non-circular) slickensides (Table 2), using the Rhinoceros curve-fitting software (version 6 SR9 6.9.18271.20591). By finding the best fit, the “ ω/v ” ratios are deduced in each cases. We could not use each and every example of curved lineations from already published figures, as in some cases the slickenlines appear to have developed on gently curvilinear surfaces, such as Fig. 3 of Tapsoba and Howard (2017).

In the first part of this study, we have analyzed lineations that have developed on planar fault surfaces. In three simple tests, the model is found to match the common sense. To cross-check whether the RStudio-derived graphs (code in Appendix 1) gave realistic results, first, when we choose $\omega/v = 0$, i.e., $\omega = 0$ and $v \neq 0$, a linear slickenside is simulated (Fig. 2a). This is the case of a purely translational fault (Fig. 1a). Second, when we take ω/v as “undefined”, i.e., $v = 0$ (such as Fig. 1b1), the slickenside appears as merely a point (Fig. 2.b). Since $v = 0$, the point does not translate to attain a linear geometry, rather it only rotates about itself. Third, higher the ω/v magnitude, more curved is the slickenside. For example, compare the curve in Fig. 2.c1 which is having less curvature than that in Fig. 2.c2, where ω/v are 1.3 and 5.6, respectively. The lineation geometries simulated in the work remain the same for normal, reverse and strike slip movement of the translational component of faulting.

The model curves produced by RStudio (sub-figures c in Fig. 3 and Rep. Figs. 3–11). The Rhinoceros curve-fitting software was used to find the correlation coefficient (R) between the traces of lineations (sub-figures d in Fig. 3 and Rep. Figs. 3–11) and The results viz., the “ ω/v ” ratio and the R values are presented against the corresponding slickenslides in Table 1.

3.2. Curved fault planes

In the second part of this work, the software Matlab (Version 9.3, 2017) was used to simulate the geometry of the lineations in 3D developed on listric fault planes that are perfectly spherical (Appendix-2). Fig. 4 presents such a lineation pattern when the hanging wall block moves with respect to the footwall block along a specific geographic direction with a constant speed (represented also by “ v ” in this work so that one can compare this with the previous case of “constant velocity”). Since the faulted block slides past along a curved surface, using the term “constant velocity” would be inappropriate. Fig. 5 represents the lineations for a perfect rotational movement of the faulted block. Fig. 6 represents the lineation geometry when $\omega = 0$. Note the lineation in this case is a segment of a circle, as expected. Rep. Figs. 12 to 22 are the several cases of lineations produced by different “ ω/v ” magnitudes. As no published photographs of lineations exist on spherical listric fault planes till date to our knowledge, correlating them with the natural examples is indeterminate.

4. Discussions and conclusions

We simulate lineation pattern on planar and curviplanar fault planes for different ratios of rates of rotation and translation (ω/v) of the faulted blocks. With the computer codes, such ratios from field photographs of fault planes containing slickenslides can be obtained and their tectonic implications can be interpreted. For example, if one ignores the rotational component of fault blocks and apply Anderson's theory of faulting to decipher the principal stress orientations, it will be just an

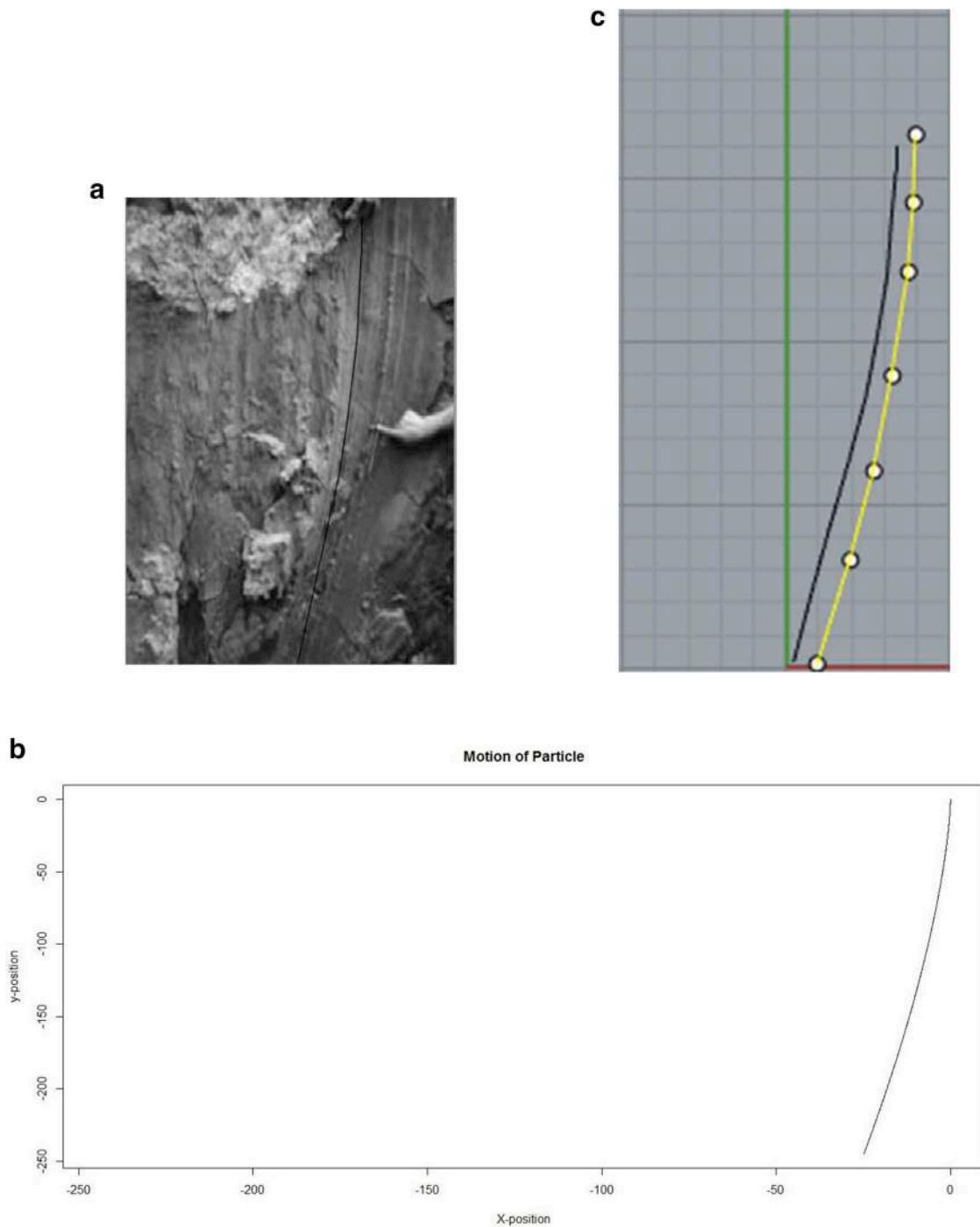


Fig. 3. **a.** Natural lineations on a fault plane from the Wenchuan region (China; Liu-Zhang et al., 2010; vide Table 1). Trace of the natural lineation is drawn. **b.** The modeled lineation for specific magnitudes of “ ω/v ” ratio, obtained by using the RStudio (Version 1.0.153-2009-20017). Here the units along X- and Y-axes represent the distances in a 1:1 scale. **c.** Fit between curves in the sub-figure **b** and the sub-figure **c**. Vide code in Appendix 1.

approximation, and will be an imperfect interpretation. Analysis of the previously published images reveals the ω/v variation from 0.2 (San Miguelito range, Central México) to 2.4 (Hyogoken-Nambu, Japan).

Different kinds of linear marks can develop on fault planes (review in Resor and Meer, 2009). Such marks picked up in seismic images, known as corrugations, may not indicate slip direction of the fault (Ferrill et al., 1999). Therefore, those marks published by previous authors were avoided in analyzing in this study.

This second part of this work considers spherical geometry of the listric fault plane. Modeling listric fault surface in terms of a well known curve or surface is commonplace in structural geological/ seismicity related analyses (e.g., Archuleta and Brune, 1975). For example, Schultz (1987) numerically model the genesis of the listric strike-slip faults by considering circular arc-shaped curved fault planes for simplistic analyses and relate curvature of fault planes and their slip amounts. Schultz (1992) describes mechanics of listric faults in 2D. For

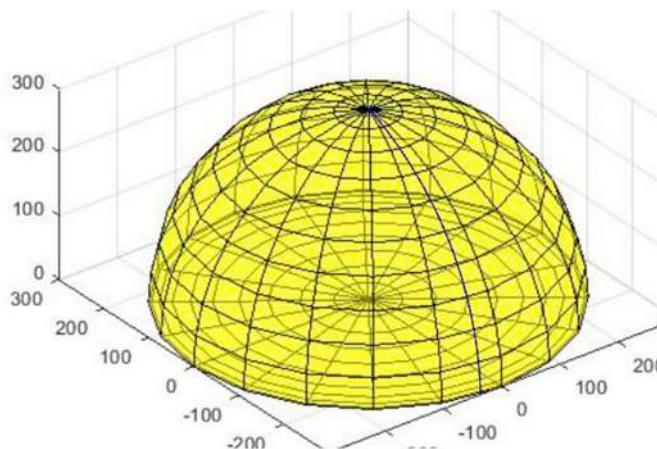


Fig. 4. Curved slickenside on a part of a spherical surface with radius = 200 units. $\omega = 1$ unit per unit time.

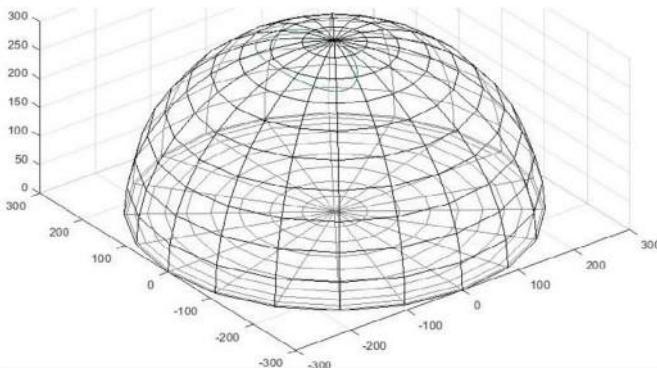


Fig. 5. Curved slickenside (blue-coloured) with perfect rotational motion on a part of a spherical surface with radius = 200 units. Radius of sphere 300 unit, $\omega = 3$ units per unit time. The X-, Y- and the Z-axes are the dimensions in three directions using which the sphere is drawn.

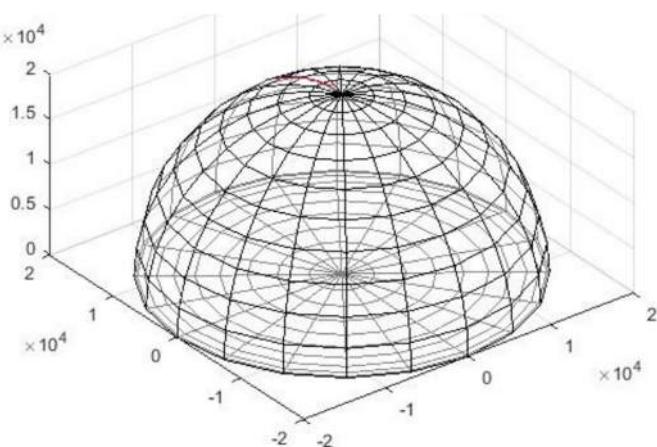


Fig. 6. The purpose of this figure is to demonstrate the lineation geometry when $\omega = 0$ on a part of a spherical surface with radius = 20000 units. We chose a radius in this case much more than that in Fig. 5 in order to observe the curved lineation at a higher resolution. Here the lineation generated (red-coloured) is the segment of a circle.

the sake of simplicity, his model considers the fault plane to be a circular arc. [Mukherjee and Agarwal \(2018\)](#) consider circular fault plane in their 2D model of listric fault-induced shear heating. [Ellis and](#)

[McClay \(1988\)](#) and [Lohr et al. \(2008\)](#) also presume cylindrical geometry of the fault plane while modeling listric fault kinematics.

However, rarely a parabolic fault plane is also considered in seismicity-related modeling ([Griffin and Davies](#), web reference). A parabolic/paraboloid slip surface sounds mechanically feasible since such a geometry of the slip surface has been deduced from Mohr-Coulomb failure criteria in slope-stability studies (e.g., [Wriggers et al., 1990](#)). If the fault plane in 3D is not perfectly spherical and the 3D coordinates of its surface are known then such set of coordinates can be input in the Matlab program. Subsequently the lineations pattern for the different magnitudes of “ ω/v ” can also be obtained.

5. Model limitations

The natural faulting process can be more complicated than the “block models” shown in [Billings \(1954\)](#), which we follow in this work. At least seven issues are there. First, in block models, the faulted blocks are rigid, therefore the nature (magnitude and direction) of net slip is uniform along the fault trend. However, in reality, (i) a regional fault can change the slip pattern along its length, from thrust slip to oblique-slip ([Gorum and Carranza, 2015](#)); and (ii) the fault slip is maximum usually at the middle portion of the fault length and decreases (uniformly) towards the two tips ([Ron et al., 1984](#); [Feng et al., 2016b, 2018a,b](#); [Fossen, 2016](#)). These mean that the faulted blocks, in a regional-scale act as deformable solids. Second, [Amos et al. \(2007\)](#) report listric reverse faults from the Mackenzie Basin, New Zealand. When listric faults form a network ([Claypool et al., 2002](#)), the kinematics is expected to be more complicated. Third, faults with rotation components in their slip have been cursorily referred to have temporally shifting pivot of rotation ([Billings, 1972](#)). Fourth, [Pueyo et al. \(2002\)](#) deduce from paleomagnetic studies from the External Sierras (Southern Pyrenees) an acceleration in rotation rate of faulted blocks, from 5.5 to 9° Ma^{-2} . A ten times temporal increase of rotation rate has also been reported by [Mochales et al. \(2012\)](#) that has been linked with the emplacement of a thrust sheet. Fifth, A few natural faults have been reported where the geometries of curved lineations on the fault planes vary (e.g., Rep. Fig. 24). This means that over the same fault plane, “ ω/v ” ratio can change. Sixth, curvature of fault planes may vary temporally ([McClay and Ellis, 1987](#)). Seventh, Landslides and glacial abrasion sometimes leave curved lineations on (curved) planes. We have not dealt fault networks, shifting pivot, accelerated rotation of fault-blocks, different “ ω/v ” ratio at different portions of the faulted block, changing curvature of fault planes, and landslide and glacier-related lineations in this work.

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Table 1

Rotational information from different terrains in the world.

Author(s)	Locations	Rotation rate ($^{\circ} \text{Ma}^{-1}$)	Total rotation (in $^{\circ}$)
Pueyo et al. (2002)	Thrust system in External Sierras (Southern Pyrenees)	5.3 to 21	At least 40
Mochales et al. (2012)	Thrust belt around Boltaña anticline, Southern Pyrenees	1 to 10	52
Rodriguez-Pinto et al. (2016)	Around Balzes anticline (Southern Pyrenees)	5.2	
Oliva-Urcia and Pueyo (2007) and ref. therein	West of South Pyrenean Central Unit	25 to 50	
Oliva-Urcia and Pueyo (2007) and ref. therein	Eastern part of South Pyrenean Central Unit	80	
Oliva-Urcia and Pueyo (2007) and ref. therein	Internal Sierras, southwestern Pyrenees	08 to 26	

Table 2

Numerical details of curve fitting exercise.

Sl. no.	Previous workers	Previous workers' fig. no.	Location (See Repository Fig. 1 for plot in the world map)	ω/v specified after curves were fitted	Fig. no. in this article	The correlation coefficient (the R-values)
1	Liu-Zang et al., (2010)	12	Hongkou (China)	0.6	Main text: 3a	0.96
2	Xu et al. (2013)	4a	San Miguelito range (México)	0.4	Repository: 3a	0.95
3	Xu et al. (2013)	4c	San Miguelito range (México)	0.2	Repository: 4a	0.95
4	Otsuki et al. (1997)	6a	Hyogoken (Japan)	1.4	Repository: 5a	0.95
5	Otsuki et al. (1997)	6b	Hyogoken (Japan)	2.4	Repository: 6a	0.95
6	Pan et al. (2014)	7c	Wenchuan (China)	0.6	Repository: 7a	0.96
7	Pan et al. (2014)	8d	Wenchuan (China)	0.4	Repository: 8a	0.96
8	Pan et al. (2014)	11c	Wenchuan (China)	0.2	Repository: 9a	0.95
9	Perrot (2013)	2.4 a	Appalachian (Quebec)	1.6	Repository: 10a	0.95
10	Singleton (2015)	4a	Metamorphic core complex, west-central Arizona (USA)	0.6	Repository: 11a	0.96

Appendix C. Supplementary data

Supplementary data related to this article can be found at <https://doi.org/10.1016/j.marpgeo.2019.04.024>.

Appendix-1

The RStudio (Version 1.0.153-2009-20017) code:

```
radial_vel <- -0.05 # in m/unit_time
ang_vel <- -0.01 # in rad/unit_time ### One can, in general,
have these as functions of t
t < - 0
x <- c(0)
y <- c(0)
steps <- 1000
step_size <- 0.01
for (i in 1:steps){
  t <- t + step_size
  x <- c(x,x[i]- 0.5*radial_vel*cos (ang_vel*t*pi/2))
  y <- c(y,y[i]-radial_vel*sin (ang_vel*t*pi/2))
}
plot (x,y,main = "Motion of Particle"
      xlab = "X-position", ylab = "y-position",
      xlim = c (min(y),0), ylim = c (min(y), max(y)),
      type = "l")
z = sqrt (300^2 - x^2 - y^2)
```

Appendix-2

The general Matlab program to produce lineations on spherical surfaces:

```
[x,y,z] = sphere (20);
R = 300;
xEast = R * x;
```

```
yNorth = R * y;
zUp = R * z;
zUp(zUp < 0) = 0;
figure ('Renderer','opengl')
surf (xEast, yNorth, zUp,'FaceColor','yellow','FaceAlpha',0.5)
axis equal
hold on
z_p = [];
y_p = [];
x = [];
plot3 (x,y_p,z_p)
```

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