



RESEARCH ARTICLE

Paleoenvironmental and Paleodepositional Reconstruction of the Permian Coals From Burhar Seam IV of Sohagpur Coalfield, India: Using Integrated Mineralogical, Geochemical and Petrographic Proxies

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ABSTRACT

Deciphering the paleoenvironment from coal samples is of great academic and industrial relevance. Coal samples from the Burhar seam IV, Sharda coalmine of the Sohagpur coalfield, were analysed using a multi-proxy approach to interpret paleoclimate and paleo-depositional conditions. The analysis incorporated several techniques, viz., Ultimate Analysis, X-ray Fluorescence (XRF), X-ray Diffraction (XRD), Fourier Transform Infrared (FT-IR) Spectroscopy and coal petrography. The Sohagpur coal deposit is a part of the Lower Gondwana coal group, which is characterised by high ash content resulting from proximate analysis. While XRD and FT-IR spectra identify the mineral components in the coal, XRF results show different oxide and elemental concentrations. Micropetrography provided insights into the maceral and mineral constituents of the samples. XRD, XRF and FT-IR results indicate a warm and humid paleoclimate favoured by intense chemical weathering, evidenced by the presence of minerals such as kaolinite. Depositional conditions were further examined through sulphur content analysis using Ultimate Analysis and elemental concentrations from the XRF. These suggest deposition in a transitional environment between oxic and suboxic conditions. Petrographic analysis also supported these findings, which indicate the coal samples were deposited in alternating oxic and anoxic moor conditions, with good tissue preservation.

1 | Introduction

Coal is used as a fossil fuel throughout the world. In India, especially, coal is not just an energy source but also a foundation for the country's social and economic development

(Equeenuddin et al. 2016). India is the second-largest coal producer and stands fifth in the reserves worldwide (IBM 2021). The heterogeneity in coal comprises both organic and inorganic constituents, which are reflected by the physical, chemical, thermal and petrographic properties of the coal

Abbreviations: Ag, alginite; Arg, argillaceous mineral matter; ASTM, American Society for Testing and Materials; BET, Brunauer–Emmett–Teller; Bt, bituminite; Cb, carbonate mineral matter; Cd, collodetrinite; Cg, corpogelinite; Cl, collotelinite; Ct, cutinite; DV, dark vitrinite; Exd, exsudatinite; Fl, fluorinite; FT-IR, Fourier transform infrared; Fu, Fusinite; G, gelinite; I, inertinite; IBM, Indian Bureau of Mines; ICCP, International Committee for Coal and Organic Petrology; It, inertinite; Lt, liptinite; Lt, litrodetrinite; Ma, macrinite; MM, mineral matter; mmf, mineral matter free; P, pollen; PV, pseudo vitrinite; Py, pyrite; REEs, rare earth elements; Rs, resinite; Sb, suberinite; Se, secretinite; SEM-EDS, scanning electron microscopy–energy dispersive spectroscopy; Sf, semifusinite; Sp, sporinite; Tl-, telinite; V, vitrinite; Vt, vitrodetrinite; XRD, x-ray diffraction; XRF, x-ray fluorescence.

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TABLE 1 | Several authors worked on different Indian terrains/basins on coal/lignite deposits, and their key findings are summarised in the table.

Petrographic indices indicate the wet forest orditions with intermittent moderate to high flooding. Petrographic indices indicate the wet forest origin of these coals. The Tacher coals were formed in wet moor with intermittent moderate to high flooding. Peat accumulation in the forest mire under varying oxic to anoxic conditions with good tissue preservation. Deposition of coal mostly under limnic to limno-telmatic conditions. Some samples favouring the deposition of peat under open marsh settings formed inertinite-rich coal. Deposition under terrestrial oxic to marine terrestrial oxic to the sub-oxic condition The areas with high inertinite concentrations show periodic dry (oxic) conditions during peat accumulation. These Lignites were generated from a bog forest under ombrotrophic to mesotrophic mire conditions. A warm temperate to subtropical climate with fresh to brackish water bodies and a decreasing oxygen-poor environment characterised the paleoclimate. Due mostly to the ongoing sinking of the early Permian Gondwana basin, this sedimentary succession was deposited in terrestrial-marginal settings with marine impact. The coals of the Makum Coalfield were deposited under a mildly oxic-to-anoxic deltaic forest to limnic swamp conditions. in was warm and humid.	Author(s)	Terrain/ coalfield/basin	Methodology used	Key finding	Proxies used organic/inorganic
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The Tacher coals were formed in wet moor with intermittent moderate to high flooding. Peat accumulation in the forest mire under varying oxic to anoxic conditions with good tissue preservation. Deposition of coal mostly under limnic to limno-telmatic conditions. Some samples favouring the deposition of peat under open marsh settings formed inertinite-rich coal. Deposition under terrestrial oxic to marine terrestrial oxic to the sub-oxic condition The areas with high inertinite concentrations show periodic dry (oxic) conditions during peat accumulation. These Lignites were generated from a bog forest under ombrotrophic to mesotrophic mire conditions. A warm temperate to subtropical climate with fresh to brackish water bodies and a decreasing oxygen-poor environment characterised the paleoclimate. Due mostly to the ongoing sinking of the early Permian Gondwana basin, this sedimentary succession was deposited in terrestrial-marginal settings with marine impact. The coals of the Makum Coalfield were deposited under a mildly oxic-to-anoxic deltaic forest to limnic swamp conditions. During peat accumulation, the paleoclimate was warm and humid.	Tiru Valley	ley	Petrography	Petrographic indices indicate the wet forest origin of these coals.	Organic
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Deposition of coal mostly under limnic to limno-telmatic conditions. Some samples favouring the deposition of peat under open marsh settings formed inertinite-rich coal. Deposition under terrestrial oxic to marine terrestrial oxic to the sub-oxic condition The areas with high inertinite concentrations show periodic dry (oxic) conditions during peat accumulation. These Lignites were generated from a bog forest under ombrotrophic to mesotrophic mire conditions. A warm temperate to subtropical climate with fresh to brackish water bodies and a decreasing oxygen-poor environment characterised the paleoclimate. Due mostly to the ongoing sinking of the early Permian Gondwana basin, this sedimentary succession was deposited in terrestrial-marginal settings with marine impact. The coals of the Makum Coalfield were deposited under a mildly oxic-to-anoxic deltaic forest to limnic swamp conditions. During peat accumulation, the paleoclimate was warm and humid.	Talcher coalfield	ield	Micropetrography	Peat accumulation in the forest mire under varying oxic to anoxic conditions with good tissue preservation.	Organic
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A warm temperate to subtropical climate with fresh to brackish water bodies and a decreasing oxygen-poor environment characterised the paleoclimate. Due mostly to the ongoing sinking of the early Permian Gondwana basin, this sedimentary succession was deposited in terrestrial-marginal settings with marine impact. The coals of the Makum Coalfield were deposited under a mildly oxic-to-anoxic deltaic forest to limnic swamp conditions. During peat accumulation, the paleoclimate was warm and humid.	Bikaner-Nagaur basin	ır	Petrography, proximate and ultimate analysis, XRF, FT-IR	The areas with high inertinite concentrations show periodic dry (oxic) conditions during peat accumulation. These Lignites were generated from a bog forest under ombrotrophic to mesotrophic mire conditions.	Organic and inorganic
The coals of the Makum Coalfield were deposited under a mildly oxic-to-anoxic deltaic forest to limnic swamp conditions. During peat accumulation, the paleoclimate was warm and humid.	Mathews et al. (2020) Rajmahal Basin	а	Gas chromatography-mass spectrometry (GC-MS), inductively coupled plasma mass spectrometry (ICP-MS), isotopic study	A warm temperate to subtropical climate with fresh to brackish water bodies and a decreasing oxygen-poor environment characterised the paleoclimate. Due mostly to the ongoing sinking of the early Permian Gondwana basin, this sedimentary succession was deposited in terrestrial-marginal settings with marine impact.	Organic and inorganic
During peat accumulation, the paleoclimate was warm and humid.	Makum Coalfield	eld	XRF, XRD, Palynofacies, Petrography	The coals of the Makum Coalfield were deposited under a mildly oxic-to-anoxic deltaic forest to limnic swamp conditions.	Organic and inorganic
	Barmer basin	ι	Petrography, proximate and ultimate analysis, XRF, XRD	During peat accumulation, the paleoclimate was warm and humid.	Organic and inorganic

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FABLE 1	

Proxies used organic/ inorganic	Organic	Organic	Organic	Inorganic	Inorganic	Organic and inorganic
Key finding	The palynoflora indicates a dominance of gymnosperms, with evidence of local to regional wildfires, suggesting that frequent disturbances in continental ecosystems characterised the paleoenvironment during the early Permian in the Barakar Formation.	A wet forest swamp origin with a telmatic source and bog zone under ombrotrophic to mesotrophic hydrological conditions was suggested by petrographic indices. The existence of a freshwater environment during peat deposition was further confirmed by the correlation between macerals and total sulphur levels.	Paleoenvironment indicates a dense forest with a warm, humid climate. Depositional conditions are characterised by fluvial settings with some marine incursions	These coals were accumulated in warm, humid, to semi-arid climates with reducing to severely reducing circumstances.	Coal deposition has been considered to have occurred mainly in a warm, humid climate with varying oxygenating circumstances.	Organic matter (OM), fungal remains and charcoal suggest an oxidising environment. Mineral complexes suggest marine to continental settings under arid to humid climates.
Methodology used	Palynological studies on fossiliferous material. Macro- charcoal analysis from the Barakar Formation.	Petrography	Morphotaxonomy, palynology and organic geochemistry	XRD, XRF, FT-IR, proximate and ultimate analyses	XRD, SEM, XRF, ICP-MS	Petrography, XRD, XRF
Terrain/ coalfield/basin	Rajmahal Basin	Singrauli coalfield	East Bokaro basin	Jaintia Hills	Talcher coalfield	Sikkim (Eastern Himalaya)
Author(s)	Murthy et al. (2023)	Gopinathan et al. (2024)	Pillai et al. (2024)	Srivastava et al. (2024)	Panda et al. (2025)	Wahengbam et al. (2025)
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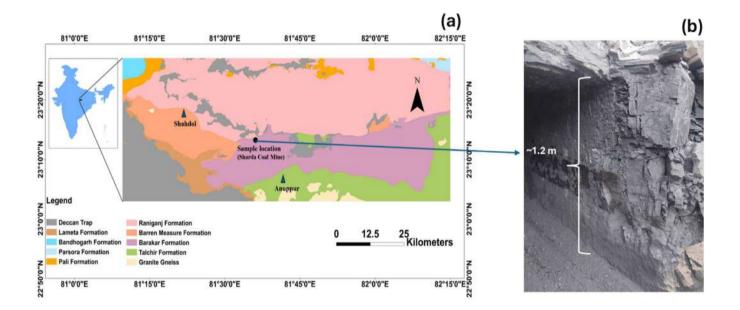
 $\textbf{TABLE 2} \quad | \quad \textbf{Geoscientific information from the Sohagpur coalfield.}$

Sl No.	Author(s)	Approach	Key findings
1	Navale and Tiwari (1967)	Palynological studies	Mostly durains constitute the coal. Comparable with Lower Gondwana units in other basins.
2	Bharadwaj and Srivastava (1970)	Spore assemblage studied from borehole coal samples	Eight coal seams were distinguished based on palynological assembelege.
3	Pareek (1987)	Microlithotypes and macerals were studied	The reflectance of coal ranges from 0.46% to 1.05%.
4	Chakrabarti (1992)	The subsidence pattern of the coalfield is studied	Characteristics of coal seam are strongly controlled by the depositional environment.
5	Ram-Awatar et al. (2004)	Palynolgical studies	Late Permian age confirmed based on palynological data, which is consistent with the middle Pali Formation in the South Rewa Basin.
6	Dhanam et al. (2013)	High-resolution seismic reflection study and Borehole sonic-log data.	The study reveals high-resolution seismic imaging of the Sohagpur Gondwana basin, indicating synsedimentary subsidence and faulting through the identification of steeply dipping normal faults and a thick Barakar Formation, highlighting the basin's tectonic and sedimentary history.
7	Agnihotri et al. (2016)	Megafossil study	Floral assemblage overall resembles that with the Barakar Formation from other Gondwana basins in India.
8	Gautam et al. (2016)	Palynological studies	Clear-cut assemblage zones of late Permian, early and late Triassic ages.
9	Jasper et al. (2017)	Scanning electron microscopy	Fire events in a peat-forming environment.
10	Agnihotri et al. (2018)	Palynological studies	Thick coal seams were formed in the region as a result of the presence of lush vegetation during the deposition of the Barakar Formation of the Sohagpur Coalfield.
11	Mondal et al. (2020)	Geophysical tools	Large-scale coal removal, surface blasting, and the existence of hard roof strata all contributed to the longwall roof's high microseismic activity.
12	Mondal et al. (2021)	Geochemical, geophysical and geomechanical analyses	Comment on reserve quality, properties of the roof, properties gas content
13	Modi et al. (2022, 2024)	Study of rare earth elements	Average contents of critical REEs: Y—51.49 ppm, Nd—79.34 ppm, Eu—2.67 ppm, Tb—1.71 ppm, Dy—9.76 ppm
14	Bhapkar et al. (2023)	X-ray scattering measurements and high-resolution imaging	Mesopore volume and surface area reduce with depth.
15	Kishor, Kumar, et al. (2024)	Brunauer–Emmett–Teller (BET), SEM, FT-IR and High-pressure CO_2 adsorption	Coal samples from the Sharda coal mine show multilayer adsorption.
16	Kishor, Kumar Srivastava, and Singh (2024)	Vitrinite reflectance was studied, and calculation for methane and adsorbed gas content was made.	The rank of Sohagpur coal is of high volatile bituminous A type.
17	Rathore et al. (2025)	Determination of ultimate analysis and proximate components, major and trace elements using ICP-MS. XRD, X-ray fluorescence (XRF), petrography, FTIR and scanning electron microscopy–energy dispersive spectroscopy (SEM-EDS)	Elemental study suggests that some elements like Co, Cr, Nb, Ni, V, Y and Zn are very higher amounts than World Clark Hard Coal (WCHC) values. As some of these elements are known for their toxicity, they can be a threat to the ecosystem and human health.
18	Kishor et al. (2025)	Low-pressure gas adsorption analysis, scanning electron microscopy and (FTIR)	Specific surface area = 4.679 to 21.974 m ² g ⁻¹ , total pore volume = 0.027 to 0.054 cm ³ g ⁻¹ . Average pore diameter = 9.834 to 26.145 nm. Pores have a high capacity to store gas.

(e.g., Ward 2002; Saikia et al. 2007; Singh et al. 2019; Kumar Srivastava et al. 2024).

Organic matter of coal is crucial to define its nature and suitability for various utilities, while inorganic matter results from coal combustion at high temperatures (Creelman and Ward 1996; Ward 2002; Singh, Singh, Singh, and Banerjee 2015). Nowadays, researchers are also focusing on the evaluation of the presence and/or extraction of Rare Earth Elements (REEs) from coal, which is an inorganic constituent for various applications (e.g., Wang et al. 2003; Fu et al. 2010; Pazand 2015; Ameh 2019; Arbuzov et al. 2019; Liu et al. 2019; Kumari et al. 2023; Banerjee et al. 2024). The sulphur present in coal also plays a vital role in its combustion. The sulphur released during combustion can become a major cause of acid rain (Nath 2021). Variation of the inorganic sulphur present in coal samples is controlled by different geologic conditions. For example, the low sulphur content in certain coals is largely attributed to the limited sulphur availability in the environment during the time of peat formation and plant growth, where a greater amount of sulphur is present due to coal overlain by the marine roof, as well as sulphur may also come from seawater sulphate in the coals (Diessel 1992).

The organic (maceral) and/or inorganic matter (mineral matter) present in coal is crucial to knowing the paleoclimate and paleodepositional conditions of the coal samples. Based on the geologic ages, Indian coals are classified into two categories: Permian and Tertiary (e.g., Singh, Singh, and Singh 2015; Srivastava et al. 2025). Previously, organic matter served as the sole means to interpret the paleoclimate and paleo-depositional environments of coal samples (e.g., Singh and Singh 1996, 2000; Singh 2016a, 2016b). More recently, researchers have started examining the inorganic components of coal (e.g., Singh et al. 2019; Srivastava et al. 2024; Panda et al. 2025). Organic matter present in the coal can determine paleoclimate and paleo-depositional conditions (e.g., Kar et al. 2022, 2025; also Paul et al. 2025). For example, excessive inertinite connotes oxic depositional conditions (Singh and Singh 1996). Similarly, the presence of different inorganic constituents also states about the paleoclimate and paleo-depositional conditions. Coal with abundant kaolinite implies a warm and humid paleoclimate that resulted from intense chemical weathering, while



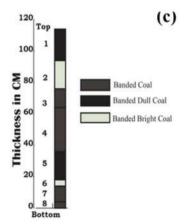


FIGURE 1 | Location map of the study area (a), Sharda mine face (Burhar Seam IV) of sample collection (b), and vertical profile section of coal samples with the thickness of each sample band (c).

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a higher concentration of illite indicates a cold and arid climate. These paleoclimatic conditions were also investigated by other elemental and oxide ratios, for example, Rb/Sr and $\rm K_2O/Al_2O_3$ (Roy and Roser 2013). The examination of both the entities (maceral and mineral matter) is crucial to constraining paleoclimate and paleodepositional conditions. Table 1 provides the details of several Indian authors who worked on the different terrain/basins (coal and/or lignite deposits) and deciphered the paleodepositional and paleoclimatic conditions using organic and/or inorganic proxies.

India possesses a huge coal reserve, which primarily occurs in the linearly distributed, isolated Gondwana basin and sporadic occurrences in the Himalayas. A limited literature survey has been done covering several isolated Gondwana basins in eastern and central India, and occurrences in the Himalaya (Table 1). The Sohagpur coalfield (Table 2) is located in the heartland of the coalproducing belts in central India, yet its paleoclimate and depositional environment were so far not been studied based on organic and inorganic proxies. Based on this research gap, the current research focuses on the detailed paleoclimatic and paleodepositional conditions based on the organic and inorganic proxies.

This study aims to decipher the paleoclimate and paleodepositional conditions of the Burhar Seam IV of Sharda coal mine from Sohagpur coalfield using multi-proxy analysis of maceral (organic) and mineral matters (inorganic).

2 | Study Area

The Sohagpur coalfield is one of the prominent coalfields, belonging to the Permian age, and is located in the southern part of the Son-Valley basin, in the central Indian state, Madhya Pradesh. In this coalfield, coal seams are mainly found in the Barakar Formation. The other stratigraphic units of the area are Raniganj, Pali, Parsora Formations and Barren Measures (Acharyya 2018). Sedimentary strata in this coalfield are ~1000 m thick and lie over

the Precambrian basement rocks (Pareek 1987; Kishor, Kumar Srivastava, et al. 2024).

3 | Methodology

3.1 | Sample Location, Collection & Preparation

For the present study, Sharda coal open-cast mine was chosen, where the coal seam associated with the Barakar Formation belongs to the lower Gondwana Supergroup (Figure 1a). In this study, 8 coal samples were collected from the top to the bottom of the mine face to maintain the total vertical profile from seam IV of Sharda coal mine, as shown in Figure 1b. The coal seam is sandwiched between the sandstones. The vertical section of the sample collection with a thickness is shown in Figure 1c.

The sample was collected using the following Indian standards, which describe the pillar sampling technique (IS:436 1964). Coal samples were kept in zip-lock bags to avoid moisture. In the laboratory, each sample has been studied according to Diesel's classification scheme for macroscopic characterisation of coal samples. The samples show banded, banded dull and banded bright natures. After that, samples were crushed and powdered in different mesh sizes for further analysis. Figure 2 details the methodology.

3.2 | Proximate and Ultimate Analyses

Proximate analysis has been carried out to determine the moisture, volatile matter, fixed carbon and ash yield present in coal samples. During the proximate analysis, -72 mesh size of powdered coal samples was used. Analysis was performed following the Indian standard (IS:1350 1984). An ultimate analysis has been conducted to determine carbon, hydrogen, nitrogen and sulphur percentages in the coal samples. Powdered coal samples

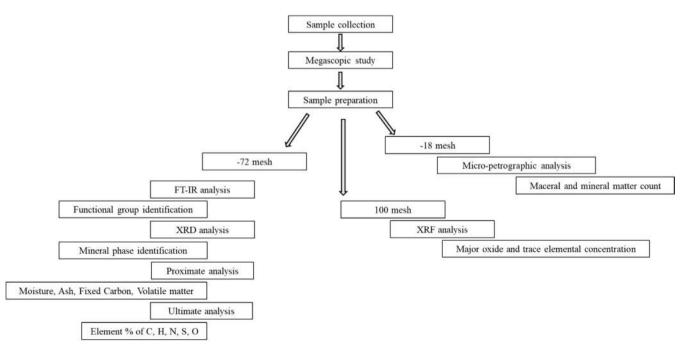


FIGURE 2 | The flowchart shows the methodology followed in this work.

of -72 mesh size were used, and a standard method of analysis was adopted (ASTM 2013).

3.3 | XRD Analysis

This preparation process was crucial, as particle size can significantly impact the results obtained from XRD, ensuring accurate identification of the various mineral phases present within the coal. The analysis was conducted over a wide angular range, specifically from 4° to 90° in the 2θ scale (θ : angle between the incident x-ray beam and the diffracted x-ray beam). This allowed a comprehensive assessment of crystal structures and mineral phases. This range was particularly effective in capturing the unique diffraction patterns that correspond to the minerals found in coal, enabling a detailed examination of its composition. For the identification of the mineral phases exhibited on the XRD chart, X'pert Highscore Plus software (v. 4.9) was used,

TABLE 3 | Coal samples and moisture, ash, volatile matter and fixed carbon % on an air-dried basis.

Sample No.	Moisture %	Ash %	Volatile matter %	Fixed carbon
1	4.76	9.12	30.00	56.12
2	5.14	13.10	25.20	56.56
3	4.03	24.20	21.60	50.17
4	3.74	22.20	23.60	50.46
5	5.46	17.50	26.50	50.54
6	9.28	13.10	27.40	50.22
7	5.17	15.10	25.80	53.93
8	4.92	22.50	23.40	49.18
Min	3.74	9.12	21.60	49.18
Max	9.28	24.20	30.00	56.56

TABLE 4 | Coal samples and C, H, N, S % on an air-dry basis.

Sample No.	C %	H %	S %	N %
1	72.02	5.25	0.90	1.56
2	69.79	4.93	0.60	1.44
3	60.69	4.12	0.52	1.15
4	62.33	4.09	0.62	1.13
5	63.42	4.66	0.85	1.31
6	64.97	5.14	0.70	1.4
7	68.54	4.64	0.52	1.26
8	61.43	4.03	0.60	1.08
Min	60.69	4.03	0.52	1.08
Max	72.02	5.25	0.90	1.56

TABLE 5 | Oxide (%) and elemental (ppm) concentration in the studied coal samples.

Samule																	
No.	SiO_2	Al_2O_3	K_2O	SO_3	TiO_2	Mn0	MgO	CaO	K_20	P_2O_5	Na_2O	Fe_2O_3	Rb	\mathbf{Sr}	$\mathbf{C}\mathbf{n}$	>	Ä
1	7.45	4.5	0.45	2.02	0.27	0.02	0.10	1.28	0.46	0.02	0.00	1.04	12.4	47.7	22.9	100.9	49.6
7	9.28	4.63	0.34	1.62	0.45	0.01	0.14	0.14	0.35	0.01	0.00	0.92	10.3	32.4	26.5	2.09	36.2
3	12.2	6.28	0.4	1.31	0.48	0.01	0.21	0.19	0.41	0.00	0.00	1.43	12.4	39.6	39.8	62	44.4
4	10.4	5.34	0.45	1.82	0.47	0.01	0.15	0.17	0.45	0.00	0.00	1.19	13	36	27.4	2.09	41.4
5	7.91	4.85	0.4	2.01	0.23	0.00	0.12	0.20	0.41	0.00	0.00	0.79	8.6	33.6	17.5	36.3	33.3
9	11.4	3.61	0.15	1.64	0.30	0.01	0.11	0.28	0.15	0.00	0.00	0.45	4.6	20	15.7	30.8	25.9
7	7.08	4.54	0.42	2.14	0.24	0.00	0.10	0.18	0.42	0.00	0.00	0.50	10.3	32.7	20.9	54.8	41.2
∞	6.91	3.26	0.41	2.03	0.54	0.01	0.09	0.20	0.42	0.00	0.00	0.61	11.8	31.9	36.7	112.1	55
Min	6.91	3.26	0.15	1.31	0.23	0.00	0.09	0.14	0.15	0.00	0.00	0.45	4.6	20	15.7	30.8	25.9
Max	12.2	6.28	0.45	2.14	0.54	0.02	0.21	1.28	0.46	0.02	0.00	1.43	13	47.7	39.8	112.1	55

which is widely employed to analyse and match the obtained diffraction patterns with a database. The software enhances the interpretive accuracy of the data, confirming the presence of specific minerals and their abundances within coal samples. The analysis provides the mineralogic characteristics of the coal, which are pivotal for determining the paleoclimatic conditions.

3.4 | XRF Analysis

X-ray fluorescence (XRF) is a highly effective and widely used analytical tool that enables researchers to determine the concentration of various elements and oxides in a range of materials. This method is particularly valuable because it

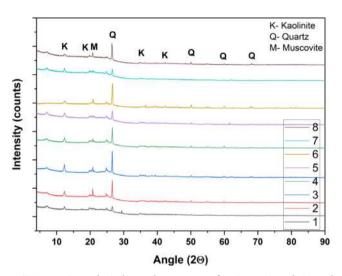


FIGURE 3 \mid Peaks indicate the presence of various minerals in coal samples. Major minerals and their peaks are presented. K—kaolinite, M—muscovite, Q—quartz.

is non-destructive, meaning that the samples can be analysed without causing damage/alteration. Powdered materials of 100 mesh size were prepared. One significant characteristic of coal is its moisture content, which acts as a natural binding agent. This property allows the preparation of pressed pellets for analysis without the need for any additional additives or chemical binders. The preparation of these pellets involves using a specialised instrument designed to compress the powdered samples. This machine applies up to 25 tons of weight to ensure that the sample packs are suitable for the analysis. Once the pellets are created, XRF analysis can be performed. The results obtained from the XRF analysis were further used to determine the paleoclimate and paleodepositional conditions of the coal samples.

3.5 | FT-IR Analysis

Fourier transform infrared (FT-IR) techniques are employed extensively to identify organic and inorganic functional groups present in coal samples. This method provides valuable insights into the chemical composition of coal, which is essential for its assessment. Coal samples were first powdered to a -72 mesh size for the FT-IR analysis and mixed with potassium bromide (KBr). KBr serves as a transparent medium that does not absorb infrared radiation. The KBr-coal mixture was pressed into a pellet to prepare the samples for analysis. The pressing process ensures that the mixture is dense and uniform, resulting in a transparent pellet that is suitable for FT-IR measurements. The FT-IR data were recorded in the transmittance mode across a 400 to 4000 cm⁻¹ spectral range. Within this range, various peaks were identified, each corresponding to specific functional groups present in the coal samples. Peaks in this range were further matched with the respective inorganic functional groups of the coal samples (Estep et al. 1968; Painter et al. 1981, 1985; Chen et al. 2015).

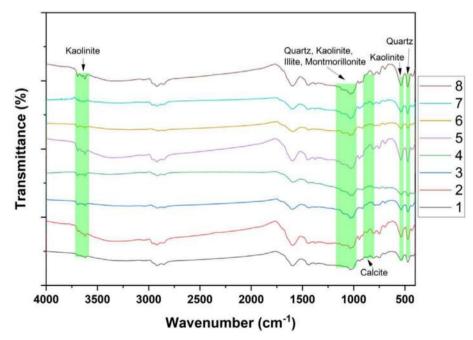


FIGURE 4 | Transmittance peaks present the inorganic functional groups present in the coal samples (shown as light green shadow).

3.6 | Micropetrography

Micropetrographic analysis was conducted on coal mounts that were explicitly prepared using -18 mesh size samples mixed with resin and hardener. This process is crucial because it allows detailed examination of the coal's microstructure, which can provide valuable insights into its properties. After the liquid mixture was poured into a mould, it solidified, forming a stable mount suitable for further analysis. The coal mount was polished to achieve a clear and scratch-free surface for microscopic examination. This involved using various grades of sandpaper and soft cloths to refine the surface until it was perfectly smooth, eliminating any imperfections that might interfere with the observations under the microscope. A reflected light microscope was employed to identify and classify macerals and mineral matter in the coal samples. The microscope was equipped with a fluorescent attachment that operated in an immersion oil medium, which enhanced visibility and contrast, enabling efficient identification. The specific model used for this analysis was a Leica microscope (DM2700P) coupled with the Leica Application Suite (Version 4.6). During the examination, macerals and mineral matter were identified using a spacing of 0.4 mm throughout the samples. This systematic approach ensures that the analysis is thorough and representative of the entire sample, enabling accurate interpretations of the coal's composition and characteristics. Throughout this process, standard procedures widely recognised in the field of coal geology were carefully adhered to in order to identify macerals under the microscope. These established protocols are critical for ensuring the reliability and validity of the results, contributing to a better understanding of coal properties and their implications in various applications (ICCP 1998, 2001; ISO 11402 2004; Pickel et al. 2017).

4 | Results

4.1 | Proximate and Ultimate Analyses

The proximate analysis result shown in Table 3 represents that the moisture content ranges from 3.74% to 9.28%, volatile matter within 21.60 and 30.00%, fixed carbon ranges from 49.18% to 56.56% and ash yield within 9.12% and 24.20% on an air-dried basis. The ultimate analysis was done to measure the sulphur present in coal samples. The C, H, N and S percent of the respective samples are shown in Table 4. Total sulphur percent is an essential parameter for deciphering the paleodepositional conditions of coal samples. The sulphur percent in these coal samples ranges from 0.52% to 0.90%.

4.2 | XRF Study

The major oxide results, ${\rm SiO_2}$ ranges from 6.91% to 12.19%, ${\rm Al_2O_3}$ ranges from 3.26% to 6.28% while K₂O ranges from 0.15% to 0.45% while Rb from 4.6 to 13 ppm, Sr from 20 to 47.7 ppm, Cu from 15.7 to 39.8 ppm, V from 30.8 to 112.1 ppm, Ga from 6.9 to 14.6 ppm and Ni from 25.9 to 55 ppm. Na₂O

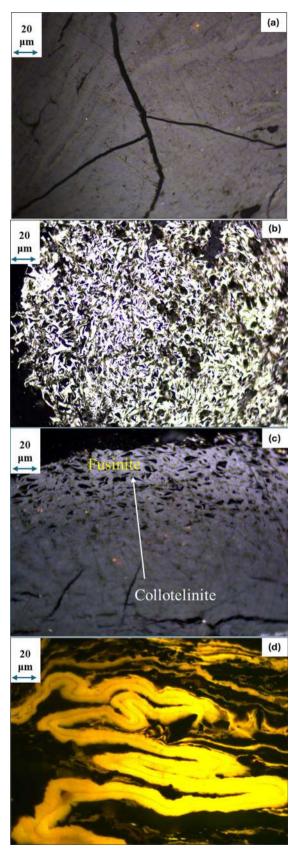


FIGURE 5 | Microphotograph of macerals present in the coal samples: a—collotelinite; b—semifusinite; c—transition from collotelinite to fusinite; d—megaspore under the fluorescent light.

deficiency in these samples is attributed to the inherently low sodium content in coal, which reflects the limited presence of sodium in the sedimentary environments where coal originates. Additionally, most sodium in coal is present in the form of soluble salts, which are often lost during the coalification process. Therefore, coal lacks significant amounts of Na $_2$ O compared to other oxides like SiO $_2$ or Al $_2$ O $_3$. Based on the result obtained through XRF analysis, the various cross-plots have been plotted to decipher the paleoenvironmental and paleoclimatic conditions. Table 5 shows the results of the oxides and elements required for the analysis.

4.3 | XRD Study

Min

0

6.1

19

0

0.3

0.3

1.2

0

2.2

6.9

28.6

1.7

12.7

There is a long history of using XRD to investigate the minerals found in coal (Dai et al. 2021). The diffracted X-ray peaks height and location in the diffractogram enable identification of the inorganic components in coal, as most of them have a crystalline

structure (Dai et al. 2021). XRD analysis was done between 4° and 90° of the angle of diffraction for the mineral's phase identification. The software X'pert Highscore plus software (Version 4.9) was used for the identification of mineral phases, and the graph was plotted on the Origin software (learning student version 2024). Various peaks present in the XRD graph of these coal samples show the presence of kaolinite, quartz and muscovite. These are the major mineral phases present in these coal samples. Figure 3 shows the mineral peaks identified in the XRD analysis.

4.4 | FT-IR Analysis

FT-IR analysis helps recognise various inorganic and organic functional groups that occur in coal samples. The analysis was performed for the wave number range between 400 and 4000 cm⁻¹ in the transmittance mode (Figure 4). The various peaks have been identified between these wave numbers. Only inorganic functional group peaks were considered in this study

TABLE 6 | Micropetrography indicates maceral and mineral matter percentages in coal samples.

Sample	Tr.I		C1	374	C 4	C-	C	DV	DX/	*7	C	C/4	D.«	۸	C1.	Т.4	T21	P4
No.	Tl		Cl	Vt	Cd	Cg	G	DV	PV	V	Sp	Ct	Rs	Ag	Sb	Lt	Fl	Exd
1	0	3	37.9	2.6	0.6	0.6	0	0	0.2	42	4.1	2.2	0.3	0	0	0.2	0	0
2	0	1	11.2	15.5	0	0	0	0	0.2	26.9	1.5	3	0.2	0	0	0.6	0	0.7
3	0.3	3	39.4	2.1	0.4	0.1	0	0	1.6	43.9	5.7	3.8	0.4	0	0	0.4	0	0.3
4	0	3	32.1	12.4	0.2	0.1	0	0	2	46.9	5.6	3	0.3	0	0	0.7	0	1.1
5	0	3	35.4	3.8	0.2	0	0	0	4.7	44.1	6.7	3.6	0.2	0	0	0.2	0	0.9
6	0	3	37.2	2.8	0.3	0	0	0	3	43.3	5.3	5	1.5	0.3	0	0.2	0.2	1.4
7	0.3	4	13.3	2	0.3	0	0.1	0	0.3	46.1	9.2	5.3	1.2	0.4	0	0.5	0	2.2
8	0	2	25.2	10.5	0.5	0	0	0	0.5	36.7	2.3	3.5	0.5	0	0	0.4	0	0.5
Min	0	1	11.2	2	0	0	0	0	0.2	26.9	1.5	2.2	0.2	0	0	0.2	0	0
Max	0.3	4	13.3	15.5	0.6	0.6	0.1	0	4.7	46.9	9.2	5.3	1.5	0.4	0	0.7	0.2	2.2
Sample																		
No.	Bt	P	L	Mi	Ma	Sf	Fu	Se	It	I	Arg	Cb	Py	MM	\mathbf{V}_{m}	nmf	L_{mmf}	I _{mmf}
1	0	0	6.8	0.2	0.5	17.3	1.7	6.1	7.9	33.8	14.2	2.4	0.8	17.3	50	8.	8.2	40.9
2	0	0	6.1	2.2	1.2	23.1	5.7	1.2	12.8	46.1	13.5	5	2.4	20.9	3.	4	7.7	58.3
3	0	0	10.6	0.2	0.8	9.6	12.7	2.5	8	33.8	9	1.9	0.7	11.7	49	0.7	12	38.3
4	0	0	10.6	0	0.3	6.9	9	3.9	7.2	27.2	13.2	1.2	1	15.3	55	5.3	13	32.1
5	0	0	11.7	0.1	1.2	9.3	5.5	8.2	9.8	34.2	7.2	1.9	1	10.1	4	9	13	38
6	0	0	13.9	0.3	0.8	8.8	3.2	2.5	15	30.7	7.5	2.8	1.8	12.1	49	0.3	16	34.9
7	0.3	0	19	0.3	0.6	7.9	3.4	4	5	21.2	8.7	3.3	1.6	13.7	53	.4	22	24.6
8	0	0	7.3	0.2	0.7	28.6	5.6	2.3	9.3	46.7	6.5	2.4	0.3	9.2	40	0.4	8.1	51.5

Abbreviations: Ag, alginite; Arg, agrillaceous mineral matter; Bt, bituminite; Cb, carbonate mineral matter; Cd, collodetrinite; Cg, corpogellinite; Cl, culinite; DV, dark vitrinite; Exd, exsudatinite; Fl, fluorenite; Fu, fusinite; G, gelinite; I, inertinite; I_{mmP} inertinite on mineral matter free basis; It, inertodetrinite; L, liptinite; I_{mmP} inertinite on mineral matter free basis; Lt, litrodetrinite; Ma, macrinite; Mi, micrinite; Mm, mineral matter; P, pollen; PV, pseudo vitrinite; Py, pyrite; Rs, resinite; Sb, suberinite; Se, secretinite; Sf, semifusnite; Sp, sporinite; Tl, telinite; V, vitrinite; V $_{mmP}$ inertinite on mineral matter free basis; Vt, vitrodetrinite.

1.2

8.2

5

15

21.2

46.7

6.5

14.2

1.2

5

0.3

2.4

9.2

20.9

34

55.3

7.7

24.6

to focus on the mineral components of the coal. The peaks within 3600 to 3700 cm⁻¹ show the occurrence of kaolinite, which falls under the inorganic functional group. Similarly, the peaks in the range of 400 to $1100 \, \text{cm}^{-1}$ come under the inorganic functional group, which includes various minerals, for example, kaolinite, illite, montmorillonite, calcite and quartz (Estep et al. 1968; Chen et al. 2015; Equeenuddin et al. 2016; Srivastava et al. 2024).

4.5 | Micropetrographic Study

Coal macerals are divided into three groups as follows: (1) Liptinite group macerals- these are comparatively rich in hydrogen, (2) vitrinite group macerals- these are relatively rich in oxygen and (3) inertinite group macerals differ chemically from the previous ones because they are substantially carbon-rich (Singh and Singh 2004). As per the International Committee for Coal and Organic Petrology (ICCP) recommendation and International Organisation for Standardisation (ISO) norms (ICCP 1998, 2001; ISO 11402 2004; Pickel et al. 2017), micropetrographic analysis was performed to identify organic (maceral) and inorganic (minerals) matters. Liptinite maceral groups were identified under the microscope attached to fluorescent irradiation (Teichmuller and Wolf 1977). Figure 5 presents the representative microphotograph of the coal macerals identified under the white and fluorescent light.

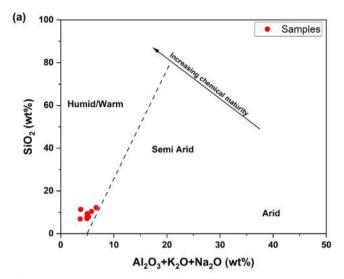
Petrographic result shows that vitrinite is a dominant maceral group, which ranges from 26.9% to 46.9%, followed by inertinite (21.2%–46.7%) and liptinite (6.1%–19.0%). The mineral matters range from 9.2% to 20.9%. Additionally, the maceral values on a mineral matter-free (mmf) basis have also been calculated (Table 6).

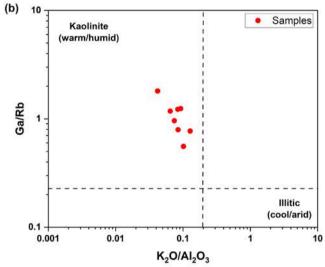
5 | Discussion

5.1 | Paleoclimatic Conditions

The rate of chemical weathering of rocks is controlled by several factors, including climate, the composition of the source rock and the tectonic setting (Mikheeva et al. 2021). A combination of XRF, XRD and FT-IR results gives an idea of the paleoclimatic conditions of coal samples. Major oxide results obtained through XRF analysis and cross-plots proposed by Suttner and Dutta (1986), followed by Mortazavi et al. (2014), Singh et al. (2019), Tchouatcha et al. (2023) and so forth show that all the samples fall within the warm/humid paleoclimate (Figure 6a). The Ga/Rb vs. K₂O/Al₂O₂ cross-plot shows that the environment was warm and humid (Figure 6b). Since samples that are rich in illite have higher values of K₂O/Al₂O₃ and lower values of Ga/Rb, while the samples rich in kaolinite show the opposite pattern with higher values of Ga/Rb and lower values of K₂O/Al₂O₃. In the case where the value of Rb/Sr is high, the climate was cold, while a lower value of this ratio represents a warmer climate (e.g., Bai et al. 2015; Jin and Zhang 2002). For the analysed samples, the Rb/Sr ratio is relatively low. This connotes a warm and humid paleoclimate (Figure 6c). The XRD and FT-IR analyses, which identify mineral phases and inorganic functional groups, suggest that kaolinite is abundantly present in these coals, since the formation of kaolinite is done by the intense chemical weathering of feldspar under

warm and humid climatic conditions; the presence of this mineral favours that coal was formed during warm and humid paleoclimatic conditions (Figures 3 and 4).





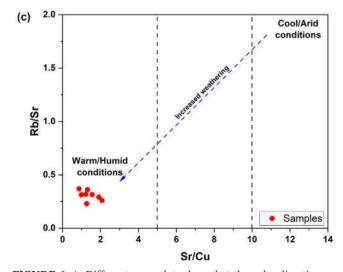
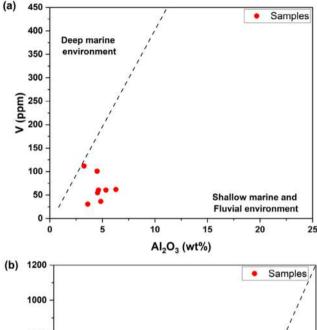
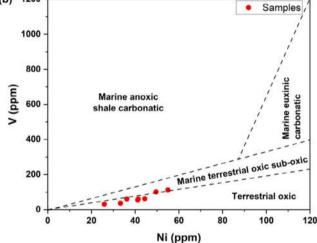


FIGURE 6 | Different cross-plots show that the paleoclimatic conditions of coal samples (a) cross-plot SiO_2 versus $(Al_2O_3 + K_2O + Na_2O)$, (b) Ga/Rb vs. K_2O/Al_2O_3 ; (c) Rb/Sr versus Sr/Cu cross-plots show the climatic conditions were warm/humid.

5.2 | Paleo-Depositional Conditions

Cross-plot of V vs. $\mathrm{Al_2O_3}$ suggests that the depositional conditions of the coal samples were predominantly influenced by environments transitioning between fluvial systems and shallow marine settings (Figure 7a). Another cross-plot explores the





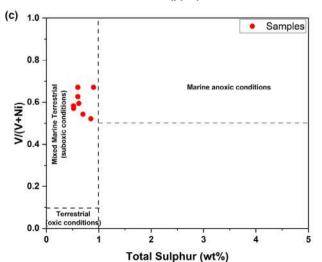


FIGURE 7 | Cross-plots showing the paleo-depositional conditions of coal samples. (a) Al_2O_3 versus V (Mortazavi et al. 2014) (b) V versus Ni (Galarraga et al. 2008), (c) total sulphur versus V/(V+Ni).

V vs. Ni relationship (Figure 7b), which serves as an essential parameter for understanding paleo deposition. The cross-plot derived from this analysis indicates that the coal samples from the Burhar Seam IV were deposited in conditions ranging from terrestrial oxic (well-aerated and suitable for organic decomposition) to marine–terrestrial oxic, suboxic environments (areas with reduced oxygen levels that can affect organic matter accumulation). Another cross-plot (Figure 7c) supports the interpretation that correlates V/(V+Ni) with the total sulphur content. The results from this plot suggest that the coal samples were deposited in suboxic conditions. This reflects a complex interplay among different depositional environments. The findings suggest that coal samples were deposited in environment with oxic to suboxic redox conditions.

Based on the micropetrographic data (Table 7), we attempt to decipher the paleodepositional conditions of the coal samples. Two depositional models were plotted as per Singh and Singh (1996) and Mukhopadhyay (1986) in Figure 8a,b.

Singh and Singh (1996) proposed a model that effectively illustrates the paleo-depositional conditions of coal samples. This model was proposed based on data obtained from the coals of the Rajmahal Basin, which is also a part of the lower Gondwana Group. The triangular plot utilises maceral and mineral matter data, which encompasses three key components: vitrinite + liptinite, Inertinite and mineral matter (Figure 8a). Vitrinite + liptinite content ranges from 33.0% to 65.1%. Inertinite values fall between 21.2% to 46.7%, while the mineral matter ranges from 9.2% to 20.9%. A comprehensive dataset of Sohagpur coals used in this triangular plot is provided in Table 7. When the maceral and mineral matter data were plotted on this triangular graph, an interesting pattern emerged—every sample point was located within the D zone, indicating alternating oxic (oxygenrich) and anoxic (oxygen-deficient) environments found in moor settings.

Maceral and sub-maceral data obtained from micropetrography were plotted in Figure 8b. This plot has been widely used by

TABLE 7 | Organic and inorganic entities (%) present in coal samples.

Sample No.	V+L	I	MM	A	В	C
1	48.8	33.8	17.3	54.4	4.2	41.4
2	33.0	46.1	20.9	20.4	20.6	59.0
3	54.5	33.8	11.7	57.6	3.1	39.3
4	57.5	27.2	15.3	50.4	16.2	33.4
5	55.7	34.2	10.1	54.6	4.7	40.6
6	57.3	30.7	12.1	59.3	3.6	37.1
7	65.1	21.2	13.7	71.4	3.0	25.6
8	44.0	46.7	9.2	35.3	12.2	52.4
Min	33.0	21.2	9.2	20.4	3.0	25.6
Max	65.1	46.7	20.9	71.4	20.6	59.0

Abbreviations: A: telinite + collotelinite + sporinite + cutinite + resinite, B: vitrodetrinite + liptodetrinite + gelocolinite, C: inertinite, I: inertinite, L: liptinite, MM: mineral matter, V: vitrinite.

several researchers for the determination of paleodepositional conditions of coal (e.g., Singh and Singh 1996, 2000; Singh 2016b; Biswas et al. 2021, 2024). The value of A ranges from 20.4% to 71.4%, B from 3.0% to 20.6% and C from 25.6% to 59.0% (Table 7). Here as well, plots fall mainly in the D zone. This indicates that coal samples deposited in a forest swamp, mildly oxic to anoxic, whereby tissues got well preserved.

From both the models (Figure 8), it is evident that the vegetal matter accumulated in the forested swamps with fluctuating water levels gave rise to alternate oxic and anoxic conditions, with good tissue preservation. The inertinite may derive from the fire in the peat (Jasper et al. 2017) or degradation of plant tissue in the oxic condition (Shi et al. 2024), whereas collotelinite

forms when the tissue preservation is good (ICCP 1998). In this case, the abundance of inertinite with a gradational change in a single grain, from collotelinite to semifusinite and finally to fusinite, was observed (Figure 5c). This might have resulted from the oxidation of the outer surface of the plant tissue, whereas the preservation of the inner or lower part of the same tissue. Hence, an alternate oxic and anoxic condition might have developed in the mire. This condition might arise due to the mixing of oxygenated water in the anoxic swamp water, which indicates a seasonal flooding of the basin.

From the petrography, it is evident that this coal is dominated by inertinite and vitrinite. Nandi et al. (1977) stated that the inert maceral is inversely proportional to combustion

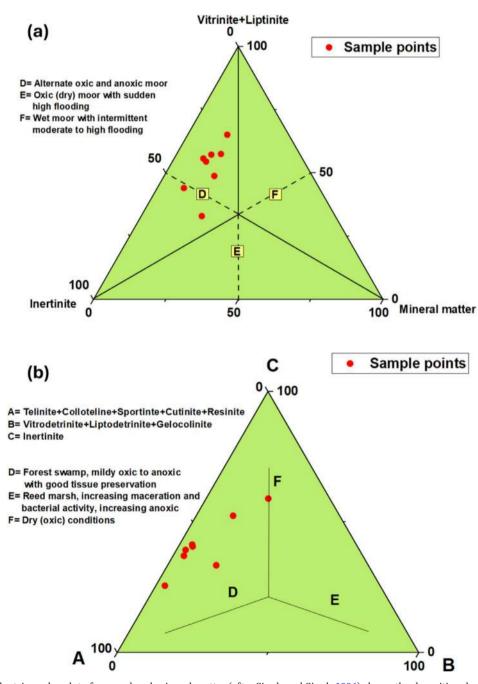


FIGURE 8 | (a) The triangular plot of maceral and mineral matter (after Singh and Singh 1996) shows the depositional conditions of coal samples. (b) The triangular plot (Mukhopadhyay 1986) based on the maceral subgroup also indicates the same depositional conditions of coal samples.

efficiency. The maceral groups also affect the CO_2 adsorption in coals; vitrinite-rich coals show higher efficiency in CO_2 adsorption due to micropore abundance (Premlall and Koech 2025). Proximate analysis shows that these coals are rich in ash yield. The inferred depositional model suggests that peat swamps suffered from seasonal flooding, which might have facilitated the influx of mineral matter into the swamp. These detrital mineral matters can be a source of REEs in coal.

6 | Conclusions

The multi-proxies and multi-analytical analyses present the paleoclimate and paleo-depositional conditions of coal samples from the Sharda coal mine. XRD, XRF and FT-IR analysis results show that during the deposition, the climatic conditions were warm and humid. Similarly, the ultimate analysis, XRF and micropetrography data were used to determine the paleo-depositional conditions. The analysis of XRF and ultimate analysis data shows that coal samples were deposited in a transitional setting of oxic to suboxic conditions. Micropetrographic analysis also shows that coal samples were deposited in alternating oxic to anoxic moor and mildly oxic to anoxic with good tissue preservation. Combining inorganic and organic proxies is essential to decipher the paleodepositional and paleoenvironmental conditions of coal samples, as Table 2 suggests that the Sohagpur coalfield has never been examined through these proxies for paleoenvironmental and paleodepositional conditions.

Author Contributions

All authors contributed to this study. K.K. collected and analysed samples and wrote the first draft. M.K.S. and G.P.T. analysed samples and revised and edited the manuscript. S.M. edited and reviewed the manuscript while A.K.S. conceptualised, supervised and finalised the manuscript.

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Conflicts of Interest

The authors declare no conflicts of interest.

Data Availability Statement

Data are provided within the manuscript.

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